**Simulating Atmospheric Freezing of Single Aqueous Droplets to Ice in a Cryogenically Cooled Ultrasonic Levitator**

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**Abstract**

Atmospheric freezing of water droplets suspended in air followed by cloud formation and precipitation represent fundamental steps of the terrestrial water cycle. These aqueous droplets exhibit distinct freezing mechanisms and thermodynamic requirements compared to bulk water often forming metastable *supercooled* water at sub-zero temperatures on the Celsius scale (< 273 K) prior to crystallization. Here, we report on a real-time spectroscopic investigation combined with simultaneous visualizations of single aqueous droplet freezing events inside a newly developed cryogenically cooled ultrasonic levitation chamber with the ultimate goal of probing the molecular structure evolution and stages of ice formation. The observed droplet freezing follows a *pseudo-heterogeneous ice nucleation* mechanism mimicking the process that occurs for atmospherically *supercooled* water droplets at the air-water interface. This proof-of-concept experimental setup allows future crystallization studies of homo- and heterogeneously doped aqueous droplets under simulated atmospheric environments - also in presence of reactive trace gases - thus untangling dynamic molecular interactions and chemical reactions, which are of fundamental interest to low-temperature atmospheric chemistry delineating with *ice nucleation* mechanisms.

The freezing of water droplets to ice is ubiquitous in nature and plays a central role in fundamental atmospheric processes such as cloud formation and precipitation including snow and hail (Figure 1). During the upward transport, water vapor condenses to droplets which continue to cool down below the freezing point of water (273 K) reaching a metastable state referred to as *supercooled water*; these undergo *ice nucleation* attaining the crystalline solid state – a pivotal step that has long piqued the interest of the scientific communities in chemistry, physics, and atmospheric sciences ([1](#_ENREF_1), [2](#_ENREF_2)). Prior experiments reveal that the freezing phenomenon largely depends on the sizes of aqueous droplets and the surrounding conditions ([3-7](#_ENREF_3)). In most of the cases, the droplet was cooled on a macroscopic surface, where the contact-effect could essentially trigger the *ice nucleation* ([8](#_ENREF_8), [9](#_ENREF_9)). Moreover, majority of the studies have been performed under vacuum for a spectroscopic characterization ([10](#_ENREF_10), [11](#_ENREF_11)). In these circumstances, fast freezing events at extremely low temperatures down to 173 K are frequently reported in the literature, which differ from the actual atmospheric situation ([12](#_ENREF_12), [13](#_ENREF_13)).

Here, we report the first combined time-dependent spectroscopic and high-speed imaging investigation of an acoustically levitated single water droplet freezing to ice inside a newly home-built cryogenically cooled process chamber thus eliminating any surface effects to reveal the temporal evolution of the molecular structure along with distinct stages of the cooling process. Atmospherically relevant conditions are simulated in the process chamber (Figure 2a, SI Appendix) starting with, e.g., tropospheric pressure ranges (760 ± 10 Torr), and as a case study for chemically inert dry nitrogen (N2) environments. The temperature of the simulated atmosphere inside the chamber and its spatial variation is monitored by a flexible silicon diode temperature sensor, while the accurate temperature of the aqueous droplet/ice particle is probed by an infrared camera (SI Appendix).

**Results and Discussion**

The water droplet is introduced to the second pressure node into the process chamber through a microneedle mounted on a wobble stick and undergoes freezing via distinct steps as the temperature drops (Figure 2b, Movie S1). For a typical 1 L sized levitated droplet, the opacity changed at a temperature of 270.9 ± 0.5 K just below the freezing point of water indicating a shift in density and initiating the formation of *ice nuclei*. However, the droplet held its fluid nature and revealed a metastable *supercooled water* until reaching the temperature of 268.9 ± 0.7 K while forming an ice crust followed by inward dendritic ice growth to the liquid core. This process represents the droplet surface induced *ice* nucleation at the interface of water and cooled nitrogen buffer gas. A small rise of droplet temperature was noted up to 269.8 ± 0.7 K due to dissipation of the latent heat from freezing. Subsequently, the droplet was rapidly transformed into a slushy ice particle, which solidified to ice at 268.5 ± 0.7 K. This process is accompanied by a slight volume increase compared to the levitated water droplet as visible. A correlation of the freezing point with droplet volume reveals a decreasing trend of the freezing point with the droplets increasing in size (Figure 2d). The relative humidity of the simulated atmosphere across the temperature range was negligible to have any significant effect on the freezing process of levitated droplet (SI Appendix).

Time-dependent Raman spectra (Figure 2e) track the temporal changes in the molecular arrangements during the freezing process by revealing significant spectral alterations. *First*, the overall band shape of the dangling O-H stretching frequency region (OH: 3800-2800 cm-1) evolved with a sharp and prominent band in the crystalline ice (hexagonal, Ih) (3144 cm-1)([14](#_ENREF_14)) which is red-shifted from the 3224 cm-1 band, i.e. the O-H stretch (OH-DDAA) of the tetrahedrally hydrogen bonded water molecule with double donor - double acceptor (DDAA) configuration in the pentameric unit ([14-16](#_ENREF_14)). This phenomenon is attributed to considerably stronger and more complex hydrogen bonding in the ice structure. However, O-H stretching vibrations of water and ice are very sensitive to their local environment with strong coupling of intramolecular and intermolecular vibrations. Therefore, spectral interpretation in terms of the normal modes, i.e. molecular symmetric and antisymmetric stretches, are fundamentally inappropriate; instead they are assigned based on localized O-H stretches. The interplay of such O-H stretches pertaining to dynamic molecular structures is evident from the gradual changes in the relative intensities of the 3224 cm-1 (OH-DDAA) and 3430 cm-1 (OH-DDA; DDA: double donor-single acceptor) bands. Both undergo spectral red-shifts by 80 and 66 cm-1, respectively, while cooling from 293.6 ± 0.9 K to 268.5 ± 0.7 K. In between 20 – 50 s, these bands, which represent *supercooled water*, are nearly equal in intensity, whereas beyond 60 s, the 3144 cm-1 band becomes predominant implying ice formation. The diminished intensity of the red-shifted OH-DDA band (3364 cm-1) in ice also indicates a structural reorganization in the bilayer framework compared to liquid water ([17](#_ENREF_17)). *Second*, a remarkable intensity reduction of the bending mode (1645 cm-1) upon crystallization is evident; this phenomenon owes to the loss of induced transition dipole moment suggesting that the concerned vibration primarily reflects water molecules that do not conform to symmetric tetrahedral environment in the liquid phase. *Third*, the appearance of structured bands at low frequency range below 370 cm-1 for ice is linked to intermolecular O…H vibrational modes, namely hindered translations of water molecules in the rigid H-bonded network.

The temporal evolution of spectrally rich OH bands upon cooling depicts that the features below the 3300 cm-1 band are enhanced, while those above 3300 cm-1 get diminished due to overall alteration in structural motifs and strengthened intermolecular interactions thus weakening O-H bonds. Hence, the aforesaid spectral segments, in particular 3200 – 3100 cm-1 (R1) and 3440 - 3340 cm-1 (R2) can be designated as characteristic features of ice and water, respectively (Figure 2f). The ratio of the integral areas under these selections enables a quantification for the progress of freezing process (Figure 2g) and avoids tedious spectral deconvolutions and consequences related to small wavenumber shifts. The resulting curve with respect to time and temperature detects the actual moment of phase change, for instance, the representative mean ratio for *supercooled water* (t = 20 - 50s) is 0.81 ± 0.07, slightly elevated than liquid water at 293 K (0.58 ± 0.04), while the bulk ice formation (dendritic growth) is triggered at ~60s as evidenced from the inflection point of the curve (1.22 ± 0.07) until reaching the Ih ice at 70s onwards (1.68 ± 0.07). The temperature-dependent temporal curve functions as a spectral marker that accurately identifies the phases and freezing point of water versus ice.

In conclusion, this study utilizing a recently developed cryogenically cooled ultrasonic levitator in a simulated environment, combined with real-time Raman spectroscopy and simultaneous visualization techniques, provides a molecular-level understanding of the freezing process for aqueous droplets by identifying the critical steps of *supercooling*, *droplet* *surface-induced* *ice nucleation at the buffer gas-water interface,* *dendritic ice growth* and *complete freezing*. This study represents the very first of its kind, where future crystallization studies will use aqueous solutions doped with atmospherically relevant liquids and solids, further mixing reactive pollutant gases with the buffer gas to identify the precise physicochemical effects on *ice nucleation* processes. Controlling the buffer gas temperature, pressure, chemical composition and constituent(s) of levitating liquid or solid sample, the apparatus can be expanded to planetary astrochemistry beyond terrestrial atmospheric chemistry.

**Materials and Methods**

The aqueous droplets were levitated at the stable pressure node in a 58 kHz ultrasonic levitator surrounded by a cooling jacket filled with liquid nitrogen. Visuals and droplet temperatures were recorded by high speed optical and infrared cameras, respectively. The temperature of the experimental environment was monitored with a silicon diode sensor. Raman spectra were obtained by exciting the sample with a 532 nm diode-pumped, Q-switched Nd:YAG laser and collecting signal via a holographic imaging spectrograph equipped with ICCD camera. The detailed experimental method may be found in SI Appendix.

**Data, Materials, and Software Availability**

All the data are included in the manuscript and/or SI Appendix.

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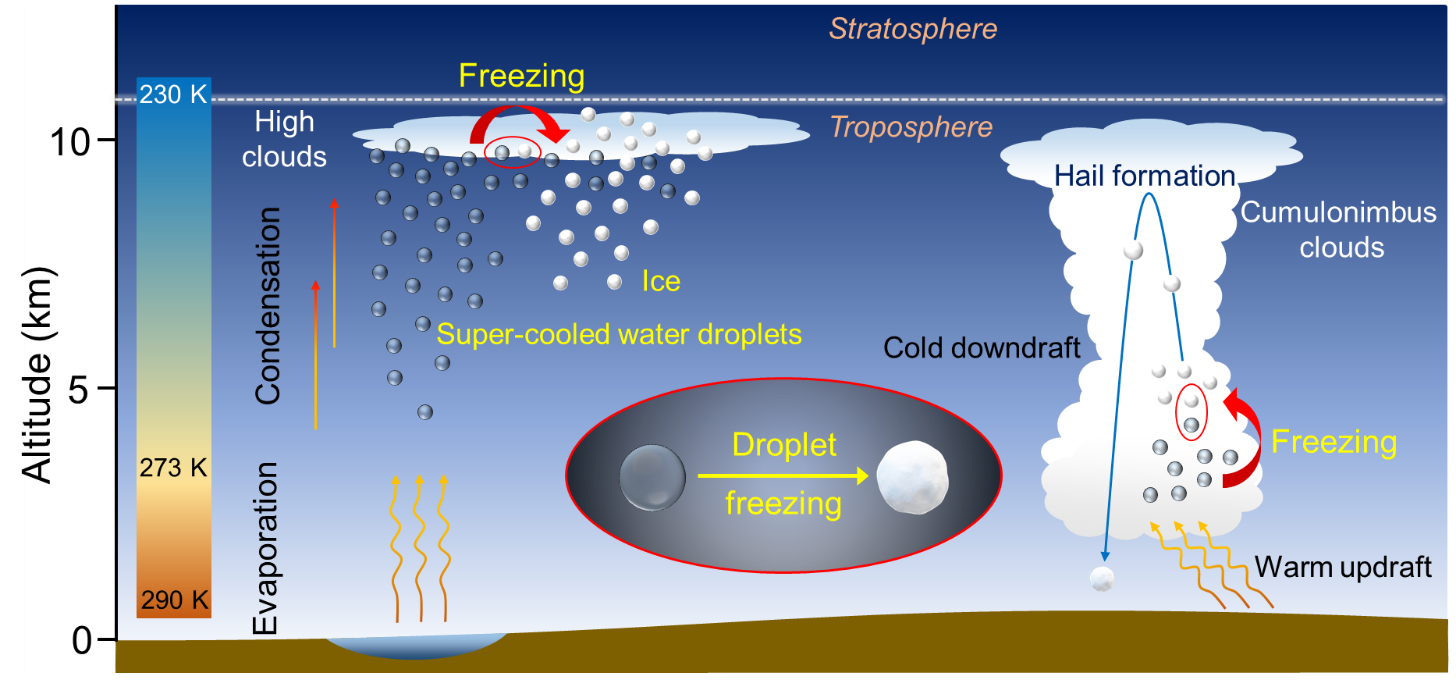
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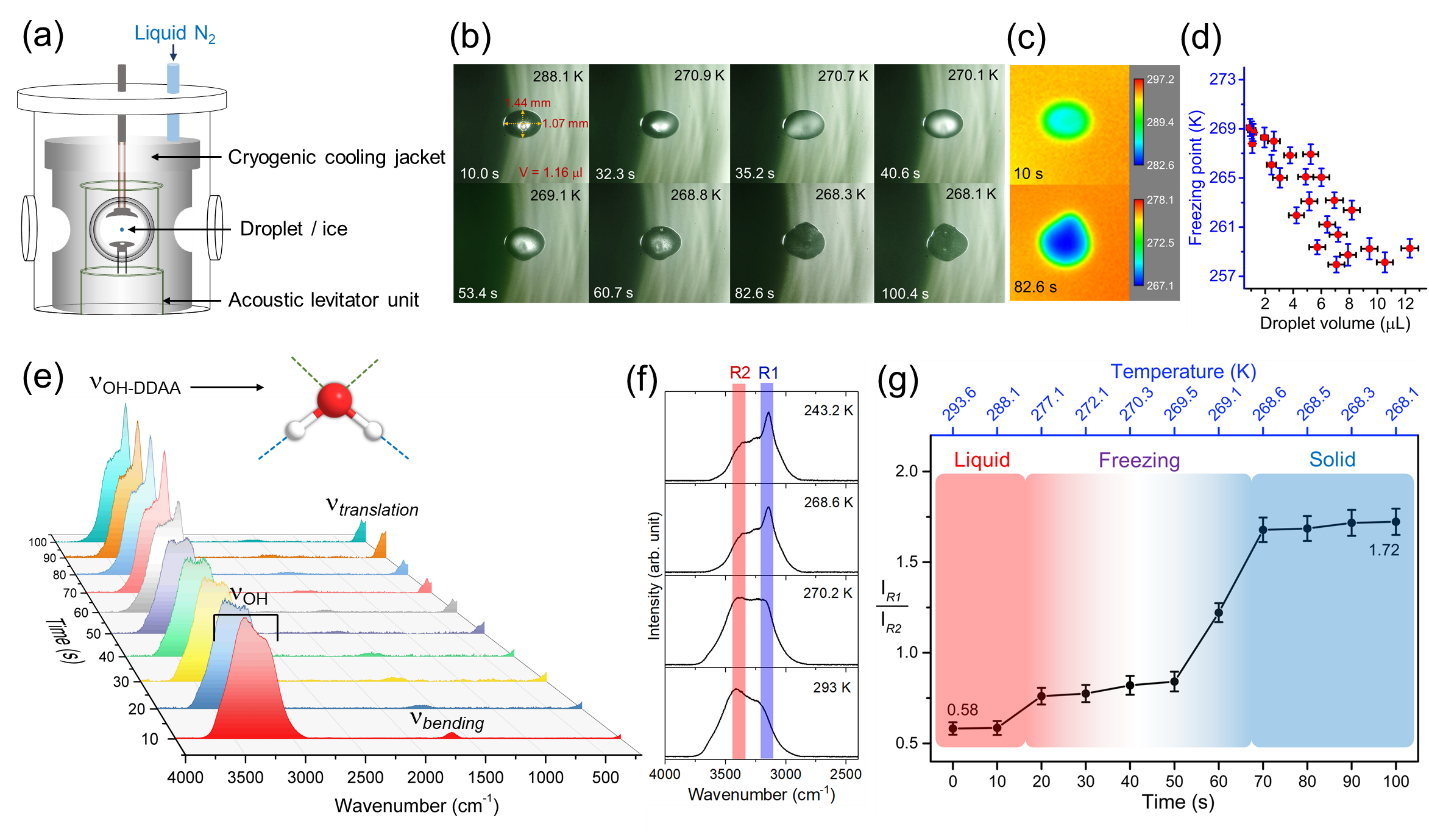
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**Figures**

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**Figure 1.** Natural atmospheric processes involving freezing of water droplets to ice as a key step.



**Figure 2.** (a) Schematic of the process chamber with a cryogenic cooling jacket surrounding the acoustic levitator unit. (b) Visualization of the freezing process of levitating water droplet versus temperature and time. (c) Representative thermal images of the water droplet and ice particle. (d) Freezing point as a function of droplet volume. (e) Time-dependent Raman spectra tracing the change in the physical state and molecular structure of water during freezing. The broad spectral region 3800 - 2800 cm-1 corresponds to the O-H stretching frequencies arising out of different kinds of intermolecular interactions. The sharp OH-DDAA band of ice represents the coupled O-H stretching of pentameric tetrahedral molecular arrangement of water with double donor-double acceptor (DDAA) hydrogen bond linkages. (f) The freezing process is quantified by evaluating the ratios of corresponding areas of two spectral segments from 3200 – 3100 cm-1 (R1) and 3440 - 3340 cm-1 (R2) representing the typical vibrational features for ice and liquid water, respectively. (g) Temporal profile of the water droplet freezing process along with droplet/ice temperature.